

# Mineralogical-petrographical observations on metamorphic transformations in the gabbroids from transitional zone of the Iuți-Tișovița-Plavișevița ophiolitic complex

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## Abstract

*On the basis of the mineralogical-petrographical observations performed on some rock samples from transitional unit of the Iuți-Tișovița-Plavișevița (ITP) ophiolitic complex in the paper are made some considerations concerning the metamorphic transformations which took place in this region.*

*The rocks were classified as chlorite-amphibole metagabbros. It has been proposed a possible succession for the identified metamorphic transformations. All these transformations had a strong metasomatic character and were determined by some hydrothermal fluids which removed from the studied rocks especially Ca, Si, and Na.*

**Key words:** *gabbroids, Iuți-Tișovița-Plavișevița ophiolitic complex, metamorphic transformations, mineralogical-petrographical observations*

## 1. Introduction

The paper presents a mineralogical-petrographical investigation performed on some rock samples taken from the transitional unit of the Iuți-Tișovița-Plavișevița (ITP) ophiolitic complex. These rock samples have been taken from the left versant of the Danube River, more precisely, from the slope (fig. 1) of the Orșova-Moldova Nouă road (DN57), uphill the bridge over Tișovița valley. This road follows the Danube River, through the ITP ophiolitic complex. The studied rock is compact and hard. It has a green color and an inequigranulare, massive isotropic texture.

The ITP ophiolitic (mafic-ultramafic) complex represents an obduction fragment of the oceanic crust, well preserved in the Upper Danubian basement [1, 2], intermediary between lithogroups Drăgșan (dominated by MT-MP metabasites) in superior position and Lainici-Păiuș (dominated by HT-LP metasediments) in inferior position [2, 3, 4, 5]. This relation must interpretation like a suture zone of the Paleozoic age, between Drăgșan and Lainici-Păiuș terranes [4]. The ITP ophiolitic complex is continue in the south in several major ophiolitic massive: Zaglavac and Deli Jovan (north-east Serbia), and Tcherni Vrach (north-west Bulgaria). In Stara Planina, is covered by a Cambrian formation of insular arch, low metamorphosed = Berkovica group [6]. It is assuming [4] to the obduction of the ITP ophiolitic complex over the rocks from Lainici-Păiuș lithogroup have Ordovician age, on the basis of the Brustur Formation age who contains cobbles

of the ITP rock-type. The oceanic crust of the ITP ophiolitic complex revealed a classical sequence of the pillow-lave, lamellar dykes, gabbroic cumulate, dunitic cumulate and harzburgitic tectonites [1, 6].

The ITP ophiolitic complex includes two main units which show the ophiolite igneous stratigraphy: a lower unit with upper mantle lithologies and an upper association of plutonic cumulates [1, 2]. The mantle peridotite unit, forming the lower part of the ophiolite igneous stratigraphy, consists mainly of harzburgite with tectonite structure and subordinate dunite, hosting small podiform chromitites. The cumulate sequence includes an ultramafic unit (dominated by layered dunites, which show lens-shaped bodies of plagioclase-bearing dunite, troctolite, olivine gabbro and gabbro), a transitional zone (consists of alternating cumulates, mafic = troctolite, olivine gabbro, gabbro, and ultramafic = dunite, plagioclase dunite, wehrlite, clinopyroxenite) and a layered mafic unit (include gabbro, olivine gabbro and troctolite, forming a rhythmic layered sequence which is cross-cut by isotropic olivine gabbro) [1, 2].

## 2. Mineralogical-petrographical investigations

### 2.1. The mineral composition

By the microscopic examination in transmitted polarized light performed on several thin sections it has been found that the rock is composed from primary and secondary minerals. Their relative proportions are variable depending on the sampling location. The primary minerals are igneous and the secondary ones are formed subsolidus. Both categories are presented in table 1.

**Table 1.** The primary and secondary minerals in the studied rock and their volume proportions.

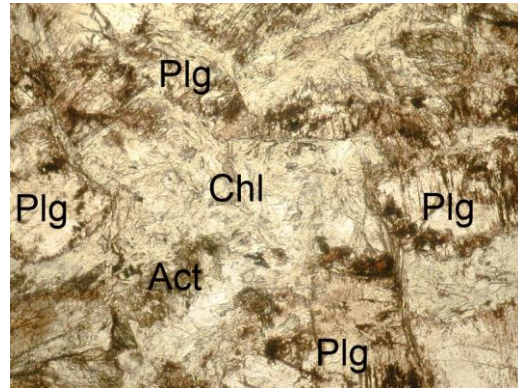
Genetic type	Mineral	Volume proportion, %
Primary minerals (igneous)	Plagioclase	35–45
	Clinopyroxene	0–5
	Magnetite	< 1
Secondary minerals (subsolidus)	Tremolite	15–30
	Actinolite	5–12
	Brown hornblende	0–2
	Chlorite	18–23
	Epidote	0–1
	Zoisite	0–1
	Calcite	0–3
	Zeolite	

**The plagioclase** is a labradorite and it is found as polysynthetically twinned crystals with relict features. The majority of these relicts are surrounded or traversed by chlorite with or without fragments of amphibole (fig. 2). The plagioclase relicts have dimensions ranging from 0,6 to 3,2mm.

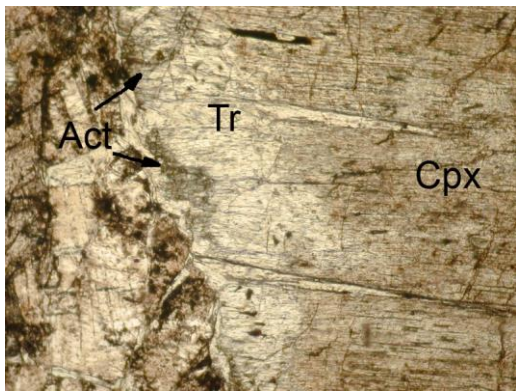
**The amphibole** is found only as pseudomorph crystals after some older crystals of clinopyroxene. These pseudomorphs are either complete ones either partial ones (fig. 3). They present a particular zonation: in the central zones the amphibole is a tremolite (colorless amphibole) and in the peripheral zones it is an actinolite (green and colorless amphibole). The transition between these two types of amphiboles is often a sharp one and seldom is gradually (fig. 4). These relations show that they are formed one on the other. Locally the actinolite is transformed in a brown amphibole by an oxidation process.



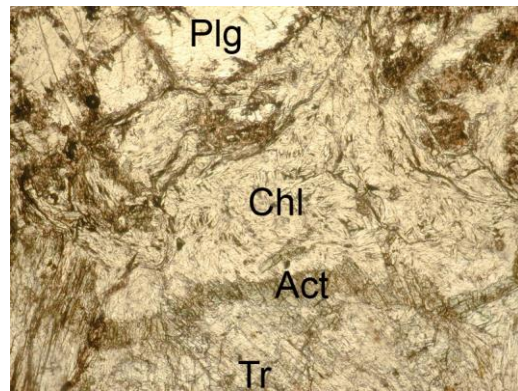
**Fig. 1.** Outcrop in the igneous rocks from slope of the Orșova-Moldova Nouă road = DN57 (uphill the bridge over Tișovița valley).



**Fig. 2.** Plagioclase relics surrounded by magnesian chlorite (N11 40x).



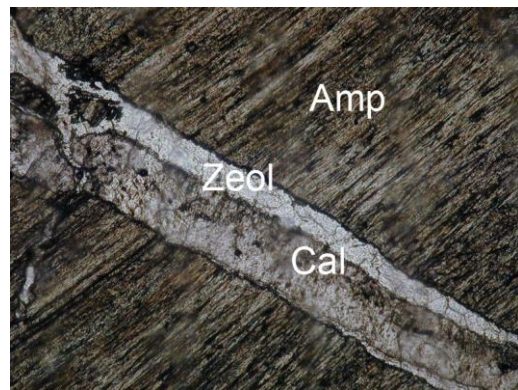
**Fig. 3.** Partial amphibole pseudomorph on clinopyroxene: actinolite at the margin; clinopyroxene in the center; tremolite between the actinolite and the clinopyroxene zones (N11 40x).



**Fig. 4.** Tremolite actinolization at the previous tremolite-plagioclase boundary (N11 40x).



**Fig. 5.** Zoizite found in the fractures in the plagioclase (N11 40x).



**Fig. 6.** Zoned joint filled with zeolite and calcite (N11 40x).

**The clinopyroxene** is found as relict nuclei surrounded through substitution by the above mentioned amphiboles. Often in these nuclei are found numerous tremolite crystals with slight discordant optical orientations giving to these pyroxene relics a perthitic aspect. In contrast with the tremolite, the clinopyroxene has a pale brown tint and numerous fine opaque inclusions.

**The chlorite** is a colorless magnesian variety. It is present as tabular crystals having dimensions of the order of 0,08–0,18mm. These crystals have random orientations. It has normal gray-white

first-order birefringence colors. Most of it is found around the plagioclase relics or in the joints which cross the plagioclase.

**The epidote** is seldom found as anhedral crystals preferentially formed in the plagioclase relics.

**The zoisite** is seldom found on the fractures which cross the plagioclase relics (fig. 5). The zoisite crystals have a columnar habit. There are not direct spatial relations between zoisite and epidote in these rocks for a relative age determination to be possible.

**The calcite and zeolite** are found filling the joints which cross both the plagioclase relics and the mafic minerals. In the analyzed thin sections these joint have thicknesses of the order of 0,1–0,2mm. Often the deposition of these minerals is zoned with the zeolite to the margins and calcite in the center (fig. 6).

**The magnetite** is present in small proportion always as subhedral crystals included in the mafic silicates. Seldom are these crystals surrounded by a thin limonitic coating.

## 2.2. Textural features

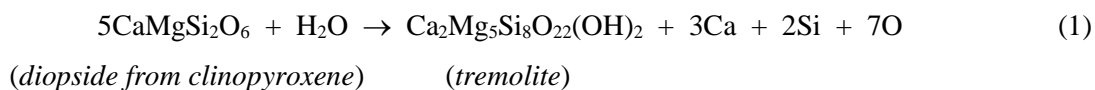
Although the substitution degree of the primary minerals is high, it is still possible to microscopically recognize the dimensions and the even the older forms of the primary crystals. Based on this kind of observations it can be deduced that the initial rock had a phaneritic structure which is specific for the plutonic rocks. A low intensity brittle deformation process led to the local fragmentation of some plagioclase and clinopyroxene crystals. The few clasts formed are also pseudomorphosed by a chloritic or amphibolic aggregate depending on the mineral nature.

## 3. Genetical remarks

The type of the primary minerals and the phaneritic structure led us to the conclusion that the rock was a gabbro *sensu stricto*. The first observed transformation that took place in subsolidus conditions was a brittle deformation. This has allowed the circulation of some hydrothermal which determined the chemical disequilibrium of the primary minerals and thus all the metamorphic transformations with a strong metasomatic character. Taking into account that the nature of the initial rock can still be identified, the fact that the transformations are metamorphic and the nature of the prevalent secondary minerals, the correct name for the studied rock is **chlorite-amphibole metagabbro**.

The most probable succession of the metamorphic transformations is:

**The pyroxene tremolitization** was a process that involves the introduction of water and partial elimination of  $\text{Ca}^{2+}$ . The reaction is:



In this tremolitization process the amphibole had a multiple nucleation on the pyroxenic paleosome. Most of the nuclei are grown epitaxial so that they the tremolite crystals have the same reticular orientation as the paleosome.

**The tremolite actinolization** is a process that consisted in cation exchange reaction between the tremolite and the metasomatic fluid. The exchange cations were:  $\text{Fe}^{2+} \leftrightarrow \text{Mg}^{2+}$ .

The reaction is:



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## Observații mineralogo-petrografice asupra transformărilor metamorfice suferite de gabbroidele din zona de tranziție a complexului ofiolitic Iuți-Tișovița-Plavișevița

### Rezumat

*Pe baza observațiilor mineralogo-petrografice efectuate asupra unor probe de roci prelevate din zona de tranziție a complexului ofiolitic Iuți-Tișovița-Plavișevița în lucrare se fac considerații genetice referitoare la transformările metamorfice suferite de rocile din zona studiată.*

*Rocile studiate au fost clasificate ca metagabbrouri cu clorit și amfibol. A fost propusă o succesiune posibilă pentru transformările metamorfice identificate. Toate aceste transformări au un puternic caracter metasomatic și au fost determinate de fluide hidrotermale care au luat în soluție din rocile studiate în special Ca, Si, și Na.*